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Novel Technique to Investigate Glacio-Fluvial Hypsometry in Hunza Using Local Indicator of Spatial Autocorrelation (LISA)

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Hypsometric Integral (HI) displays the effect of active tectonics and sensitivity on geomorphic structures. In this study we calculated HI values for Hunza valley to investigate neotectonics, development of topographic structures and process of erosion using SRTM DEM 90m. ArcGIS and MATLAB is used to generate HI and hypsometric curve (HC). We generated HI and HC values by using D8 algorithm in MATLAB to extract drainage basins for 5 and 6 Strahler orders. HI and HC values show the stages of erosion for instance high values of HI and convex HC displays young and tectonically active stage. We used different grid sizes in ArcGIS to calculate maximum, mean and minimum elevation utilizing different statistical techniques. We used Local Indicator of Spatial Autocorrelation (LISA) instead of Global Moran Index to determine the extent of distribution of clustered, dispersed and randomized HI values. This technique indicates high positive z score for auto correlated data. Regions with high HI value indicate relative uplift, undissected and young structures while low HI values indicate sediment accumulation and shallow earthquakes.

Keywords: SRTM DEM 90m, Global Moran Index, Hypsometry, Spatial auto-correlation

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Author's Contribution

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Conflict of interest

The authors declare no conflict of interest in publishing this manuscript in IJIST.

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Introduction

Hunza valley is a mountainous region situated in the northern zone of Gilgit Baltistan. The Precambrian aged rock of Hunza comprises of metamorphic and igneous rocks. The rocks in Hunza are identified on the basis of age. Hunza also comprises of the corundum [1]. Hunza comprises of two major tectonic features including Main Karakoram Thrust and Main Mantle Thrust.

Nearly 36000 km² region of MKT and MMT covers almost plutonic, subordinate sedimentary, volcanic, metamorphosed and deformed rocks. Both tectonic regions display an arc of intra oceanic island developed from the subduction of northward oceanic lithosphere [2]. The study of active tectonics is considered the most advanced tool used for the analysis of hazards and populated areas.

The geography of earth comprises of active tectonics, geographic structures and climate [3]. The geomorphology has emerged along multiple disciplines including geomorphology, geodesy, stratigraphy, geochronology, and seismology. The remote sensing techniques are used in these geomorphologic studies [4].

Hypsometric Integral (HI) displays the effect of active tectonics and sensitivity on geomorphic structures [5]. In this study we computed HI values for Hunza valley to investigate neotectonics, development of topographic structures and process of erosion using SRTM DEM 90m. ArcGIS and MATLAB is used to generate HI and Hypsometric Curve (HC) [6].

The ratio between cross section area and elevation is given by Hypsometry [7]. The area under hypsometric curve provides the value of hypsometric curve. Hypsometrically obtained data determines steps and processes of reshaping of different structures with respect to time within the defined area or natural drainage basins [8,9].

The hypsometric analysis is vital indicator of condition of watershed. It has two components including “*Hypsometric Curve*” and “*Hypsometric Index*”. HI has strong association with basin geometry, drainage area and elevation of the basin [10]. HI is independent of drainage area while it is associated with relative rate of uplift. It is scale dependant tool which may be affected by climate, lithology and tectonics [11].

Variations in glaciated terrain depend upon hypsometry, the equilibrium line altitude (ELA) moves due to changing climate [12]. According to the geometry of glacier the accumulation area and ablation of glacier change due to changing elevation. Contrariwise only a small area at the ELA being large in its height may have little variation [13].

Moreover, glaciers having similar ELA exhibit quite different behavior. One glacier with broad accumulation area and a narrow snout while other glacier with narrow upper basin and a wide tongue [15]. The percentage of area of ablation and accumulation zone of glacier can be computed by plotting ELA on HC. It also calculates the variations with respect to time [16].

SRTM DEM 90m was used to compute hypsometric integral (HI). The LISA technique is also used to calculate low and high values of HI clusters [17]. The process of erosion, climatic variations, active tectonics and lithological features of basin are determined by the shape of curve in HC.

Study Area

Hunza is a mountainous region located in the northern zone of Gilgit Baltistan. The Hunza basin is situated in the central region of Karakoram. The basin lies within the 35° 54' to 37° 05' latitude and 74° 02' to 73° 03 longitude respectively. The basin comprises of Gilgit River in the western side, China and Afghanistan in the northern zone, and Indus River and Shigar River in the southern zone. Naltar, Nagar, Shimshal and Chapursan are the major valley of the basin (Figure 1). The basin covers 16,389.4 km² out of which Hunza River comprises of 87.5% flow [18].

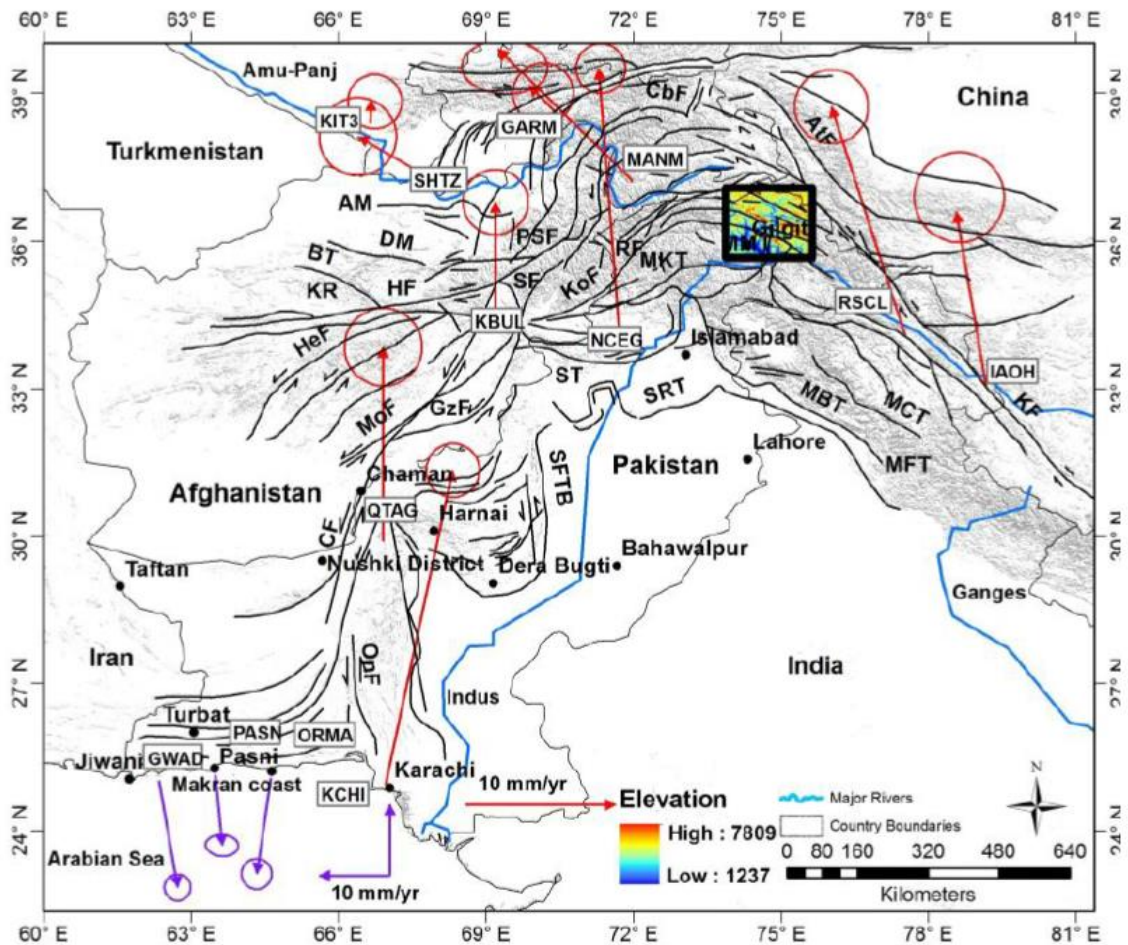


Figure 1. Diagram showing regional tectonics in the context of Indo-Pak-Eurasian collision.

Materials And Methods

Different grid sizes are required in SRTM DEM 90m to determine HI. SRTM DEM 90m is obtained from <http://srtm.csi.cgiar.org/SELECTION/inputCoord.asp>. The DEM is precise technique which covers almost all regions of the world. Nearly 16m vertical error occurs in DEM. ArcGIS fills the DEM sinks. Zonal statistics calculates the values of HI for different square grids [19]. It also computes minimum, maximum and mean value of elevations in DEM (Figure 2, 3). The obtained values of HI indicate variations of grid squares with respect to glacial, tectonic and lithological boundaries (Figure 4,5). Due to complexity of landscape both low and high values of HI can appear on the basis of grid squares [20].

Moran Index

In order to obtain the spatial distribution of HI we apply spatial autocorrelation. The spatial autocorrelation and its association are determined in geologically referenced datasets using spatial statistical analysis [21]. Figure 9. Shows the schematic illustration of cells with different HI values due to change in altitude variations depending only on the location of the cells.

$$HI = \frac{H_{mean} - H_{min}}{E - H_{min}} \quad (1)$$

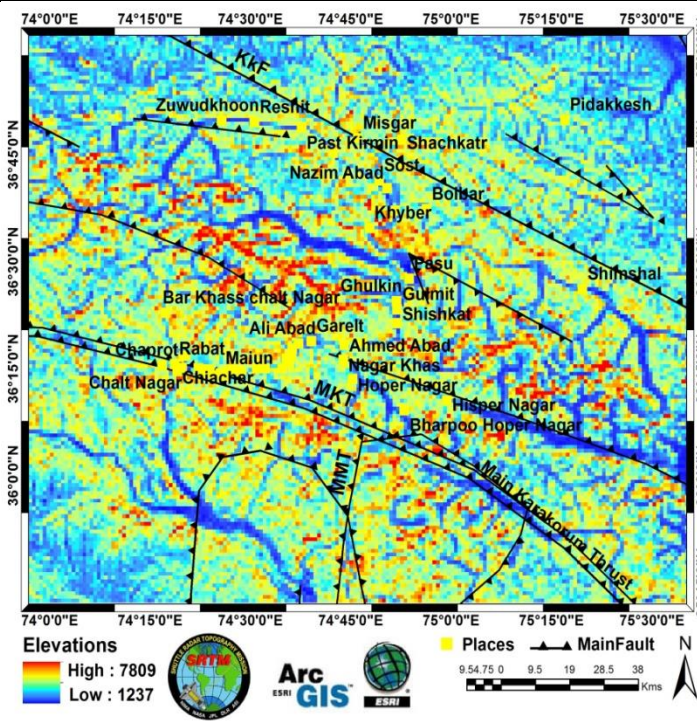


Figure.2 Digital Elevation Model (DEM) of 90 m resolution of the Hunza and adjacent regions showing faults and main rivers.

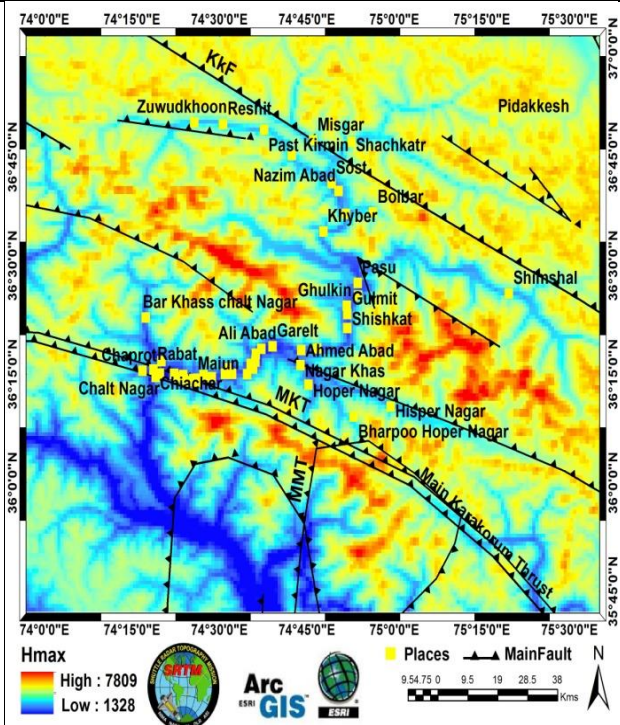


Figure. 3 Map showing spatial distribution of HI maximum values for the Hunza and its outskirts.

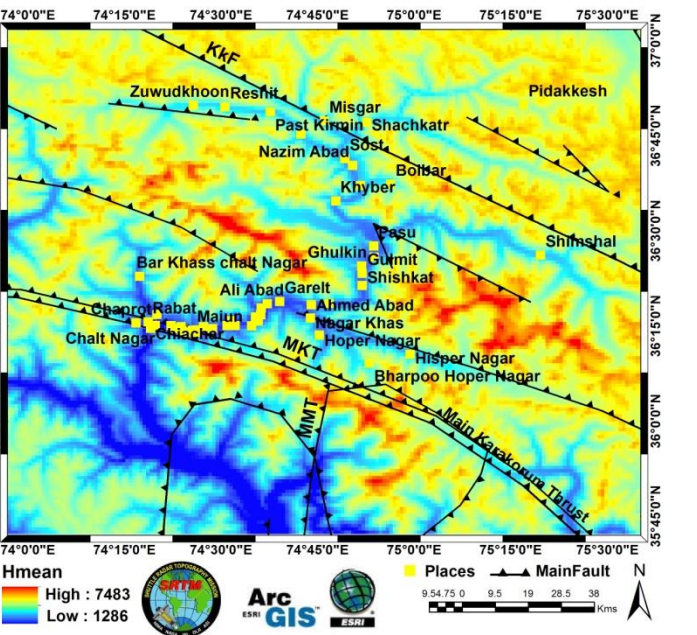


Figure.4 Map showing spatial distribution of HI mean values for Hunza and its outskirts.

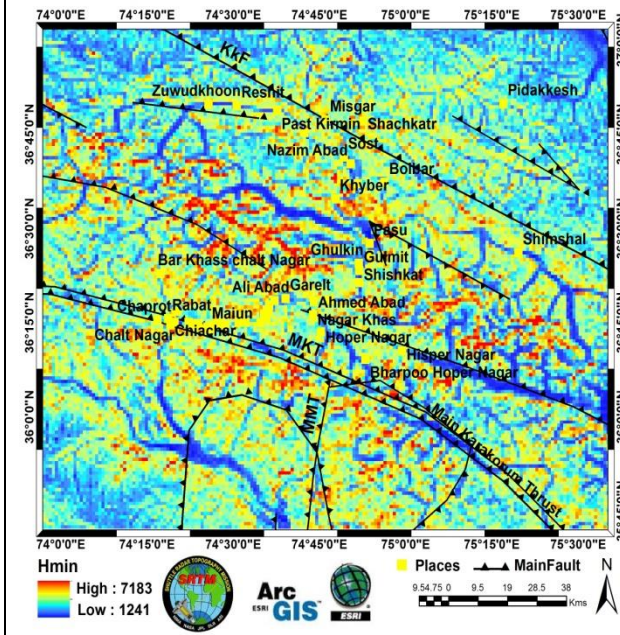


Figure. 5 Map showing spatial distribution of HI minimum values for Hunza and its outskirts.

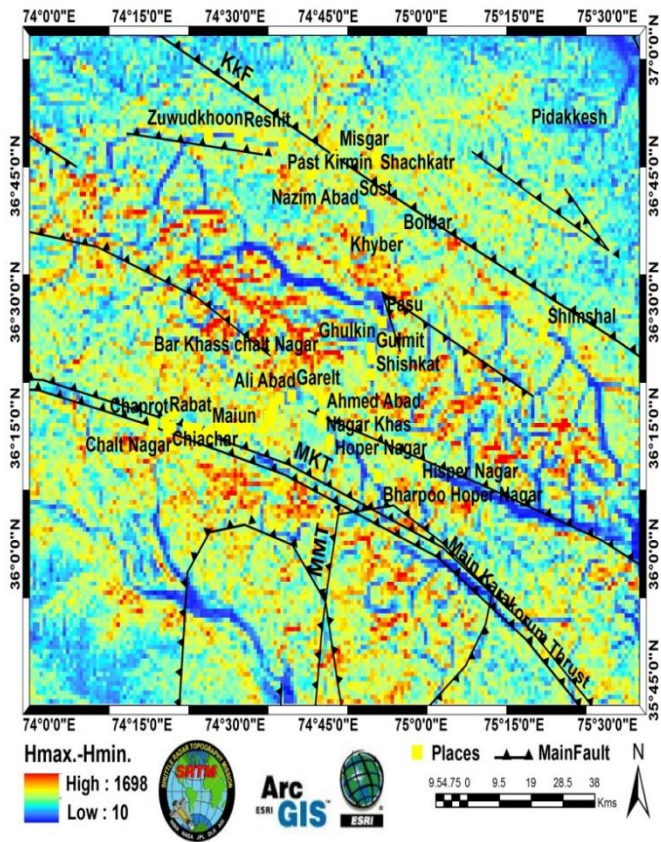


Figure. 6 Map showing spatial distribution of Hmax-Hmin values for Hunza and its outskirts.

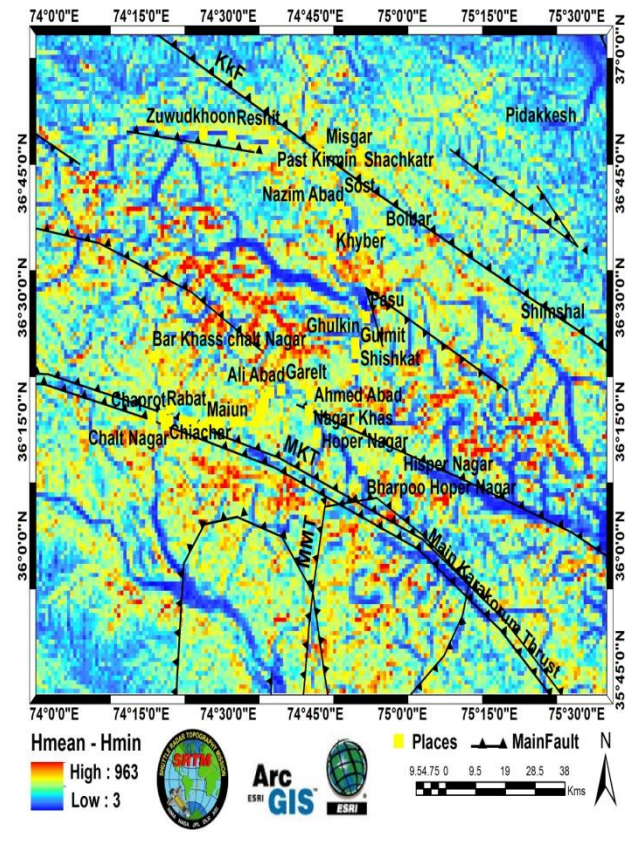


Figure. 7 Map showing spatial distribution of Hmean-Hmin values for Hunza And its outskirts.

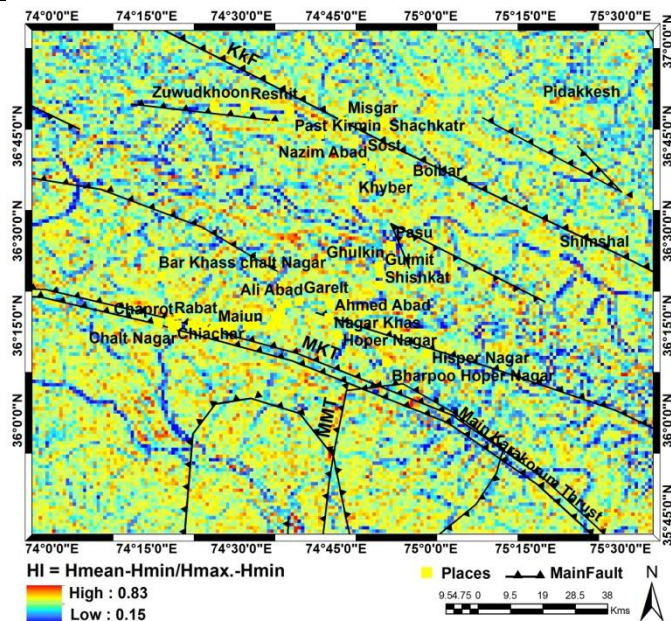


Figure.8 Map showing $HI = \frac{H_{mean} - H_{min}}{H_{max} - H_{min}}$ values for Hunza and its outskirts.

Moran I index is an accurate tool which computes the spatial autocorrelation precisely. This tool determines the dispersed, random and clustered cell values. This method correlates the observations using following formula.

$$(2) \quad I = \left[\frac{N}{\sum_i \sum_j w_{ij}} \right] \left[\frac{\sum_i \sum_j w_{ij} (x_i - \bar{x})(x_j - \bar{x})}{\sum_i (x_i - \bar{x})^2} \right]$$

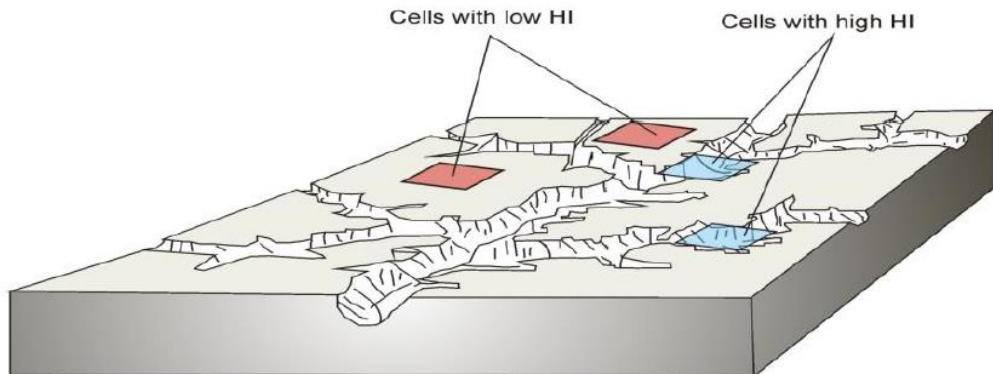


Figure 9. Schematic illustration showing cells with different HI values due to change in altitude variations depending only on the location of the cells.

Where, N represents the number of observations, i, j shows the rows and column of the matrix, w_{ij} shows the spatial weighted matrix (SWM) for each observation cells (i, j) , and x_i and x_j are the observations for the i and j location (with mean).

LISA application

GMI method provides average variation and limited association of spatial autocorrelation. LISA technique overcomes this limitation. This technique displays the spatial independence data [22]. This tool is useful for the observation of spatial clustering of global cold and hot spots at different locations. The Moran index is standard tool to determine autocorrelation while LISA explores the association among observations [23]. Specific spatial autocorrelation can be measured using the following equation.

$$(4) \quad G_i^* = \frac{\sum_j w_{ij}(d)x_j}{\sum_j x_j}$$

Where, $W_{ij}(d)$ is a weight index of i location with respect to j .

Results and Discussion

Hypsometric Analysis Basin wise

Hypsometric curve and hypsometric integral are computed for Strahler order 5 and 6 in MATLAB for Hunza sub-basins. The obtained values of HC are s shaped, concave in downward direction and convex in upward direction. The Hypsometric curve indicates different stages of erosion, active tectonics and relative rate of uplift. The high HI values are characterized by convex up of HC. It also characterizes the less corrosive regions and high topography which is under development.

Low HI values are indicated by downward concave HC (Figure 11). It also portrays erosion, old stages and even dissection of drainage sub basins. The HI values are normally distributed up to an average mean of 0.25. Distribution of HI does not display clear spatial pattern because of variations in local topography (Figure 10).

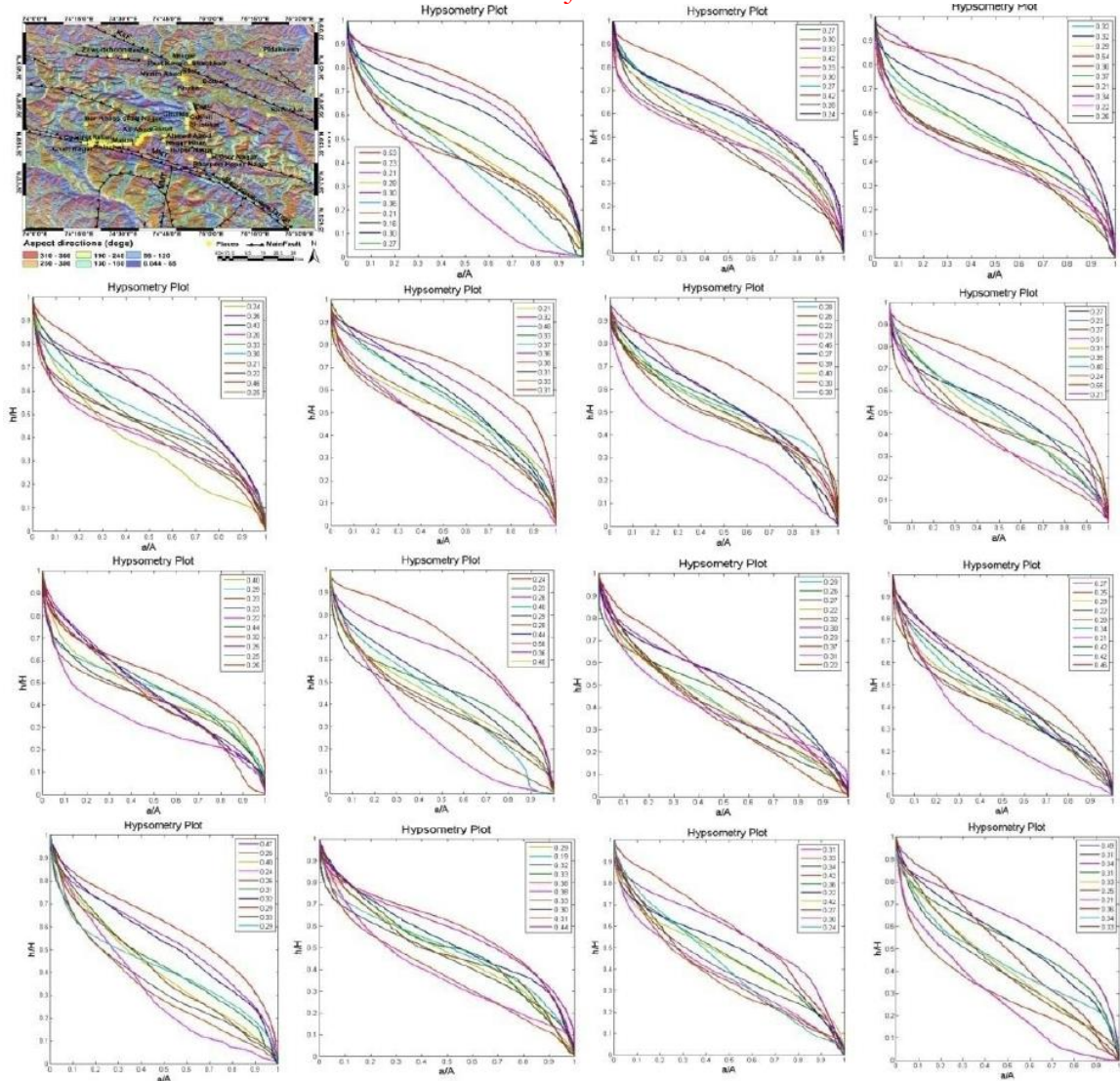


Figure 10. Illustration showing aspect map and hypsometric curves for the subbasin of order 4. **Hypsometric Analysis grid wise**

The Moran Index statistics detects the distribution of HI spatial pattern of autocorrelation Table 1 represents Moran Index, Expected Moran Index and z score for each dataset. Figure 12. shows Getis ord G^* statistics for various analysis grid sizes (1Km, 2Km and 3 Km) for a fix band distance (FBD) of 3Km.

Table 2 Results of Moran Index for 90m DEM with different analysis square grid

DEMs	Analysis Grid	N	I	$E(I)$	Z	P
90 m	1 km	27267	0.178190	-0.000040	53.624647	0.000000
90 m	2 km	6808	0.184567	-0.000159	27.730756	0.000000
90 m	3 km	3100	0.220772	-0.000346	16.558362	0.000000
90 m	4 km	1739	0.190144	-0.000614	10.515920	0.000000

High value of z-score shows the clustering of the data

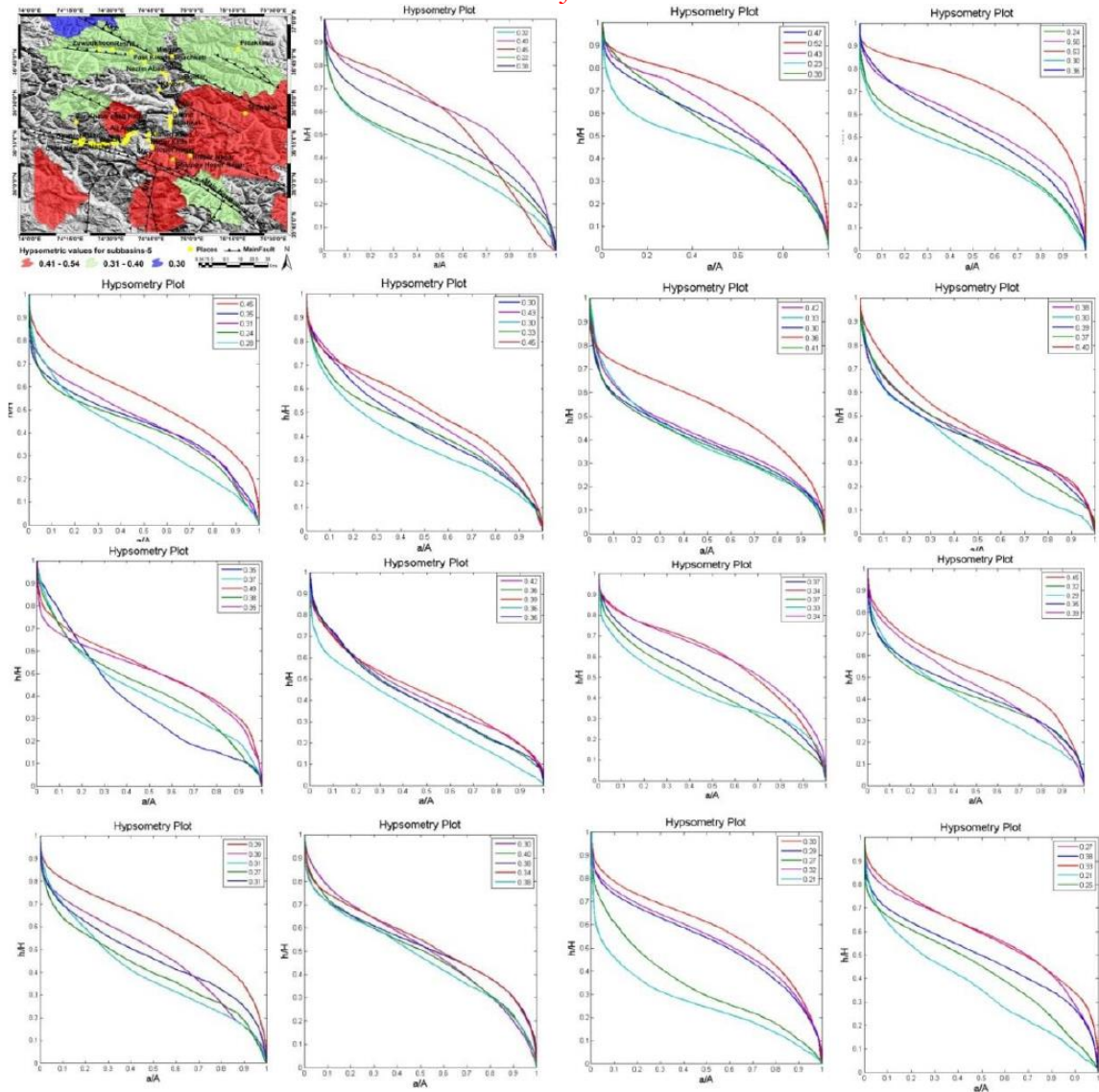


Figure 11. Illustration showing map and hypsometric curves for the subbasin of order 5.

Discussion

Spatial Dependency of scale

The global Moran I statistics display autocorrelation of HI observations.. HI datasets comprises of high positive values of z scores. The resolution of DEM decreases at different spots by increasing the grid size. The LISA technique is applied using different grid sizes; grids covered different areas although the display remained the same. The value of Z score is inversely proportional to the grid size i.e, it decreases by increasing the size of grid. The Z score exhibits scale dependency of spatial patterns of auto correlation.

Neotectonics and correlation to HI values

The result shows that the red zones and the high values of HI are distributed spatially towards northern east and southern west region of Hunza. They are most significantly distributed along northern east along MKT and MMT faults where relative rate of uplift is high. These faults are has hosted a number of intermediate, shallow and deep earthquakes. The regions with low values of HI are spatially distributed around Gulkin, Passu, Nazim Abad, Bolbar, Sost, Khyber, Shiskat, Misgar, Gulmit, Pidakkesh and Shachkatr which have sediment deposits, shallow earthquakes, fault traces and tectonically active areas.

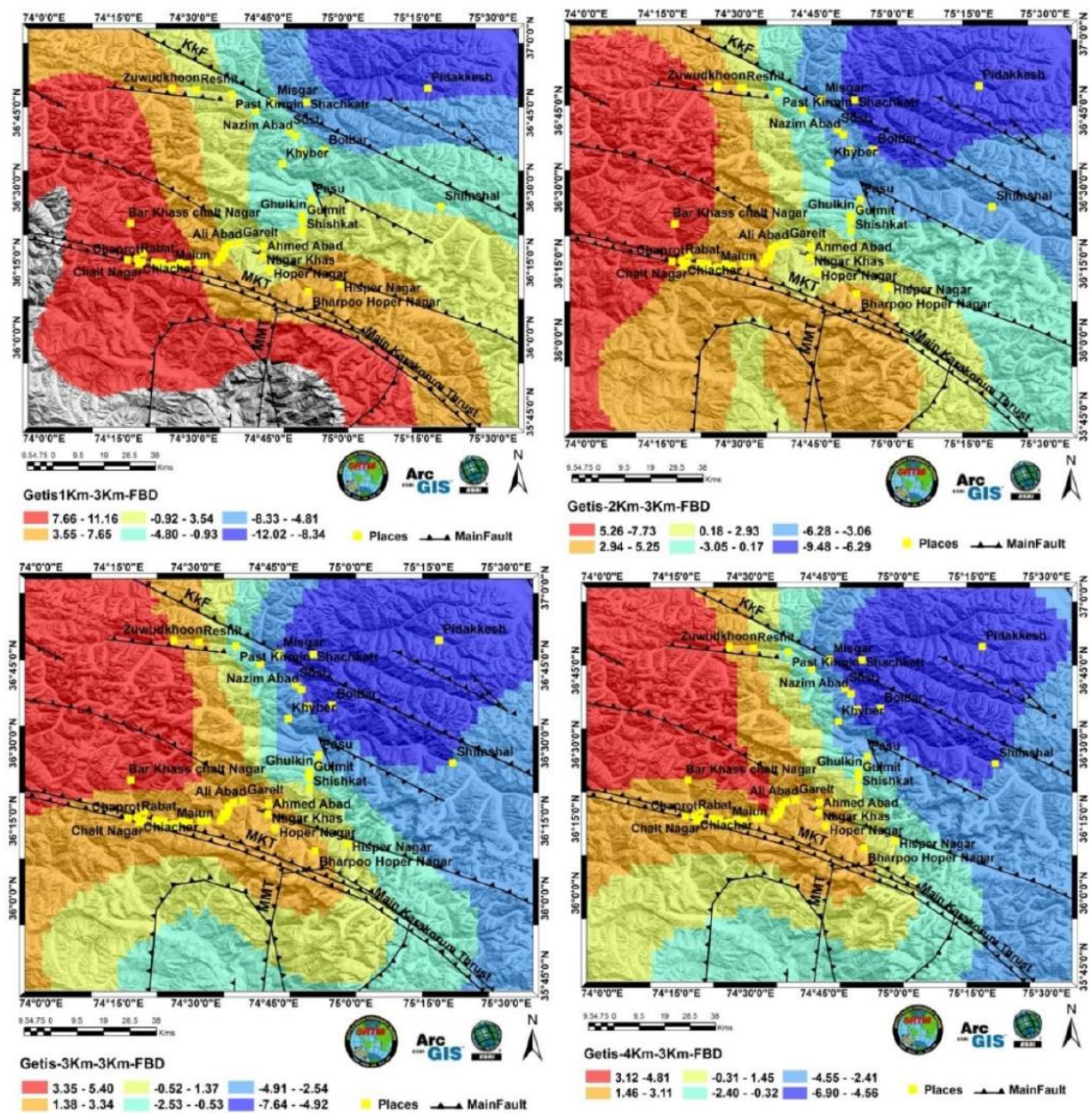


Figure 12. Map showing Getis ord G^* statistics for various analysis grid sizes (1Km, 2Km and 3 Km) for a fix band distance (FBD) of 3Km

Conclusions

Hunza valley is a mountainous region situated in the northern zone of Gilgit Baltistan. Analysis and techniques were applied over Hunza valley to illustrate that Global Moran Index (GMI) is a precise tool to compute and locate the spatial patterns in SRTM DEM 90m based datasets. It is observed from the results of GMI that HI values are distributed spatially in the form of low and high values. GMI method provides average variation and limited association of spatial autocorrelation. LISA technique overcomes this limitation. This technique displays the spatial independence data. Both these techniques illustrate that values of HI depend upon the scale. The high HI value indicates active faults and active tectonics while low value of HI indicates shallow earthquakes and sediment deposition.

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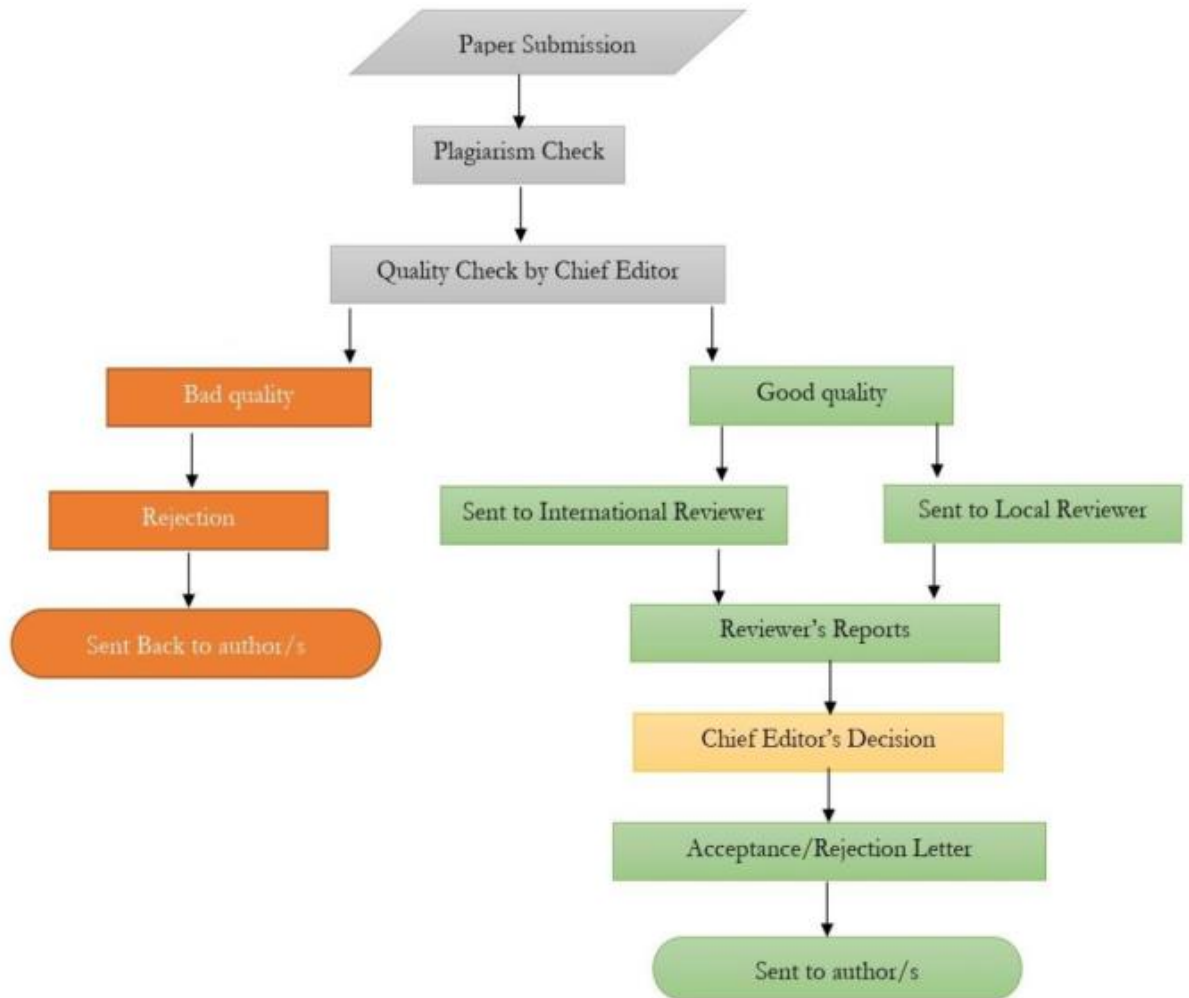


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